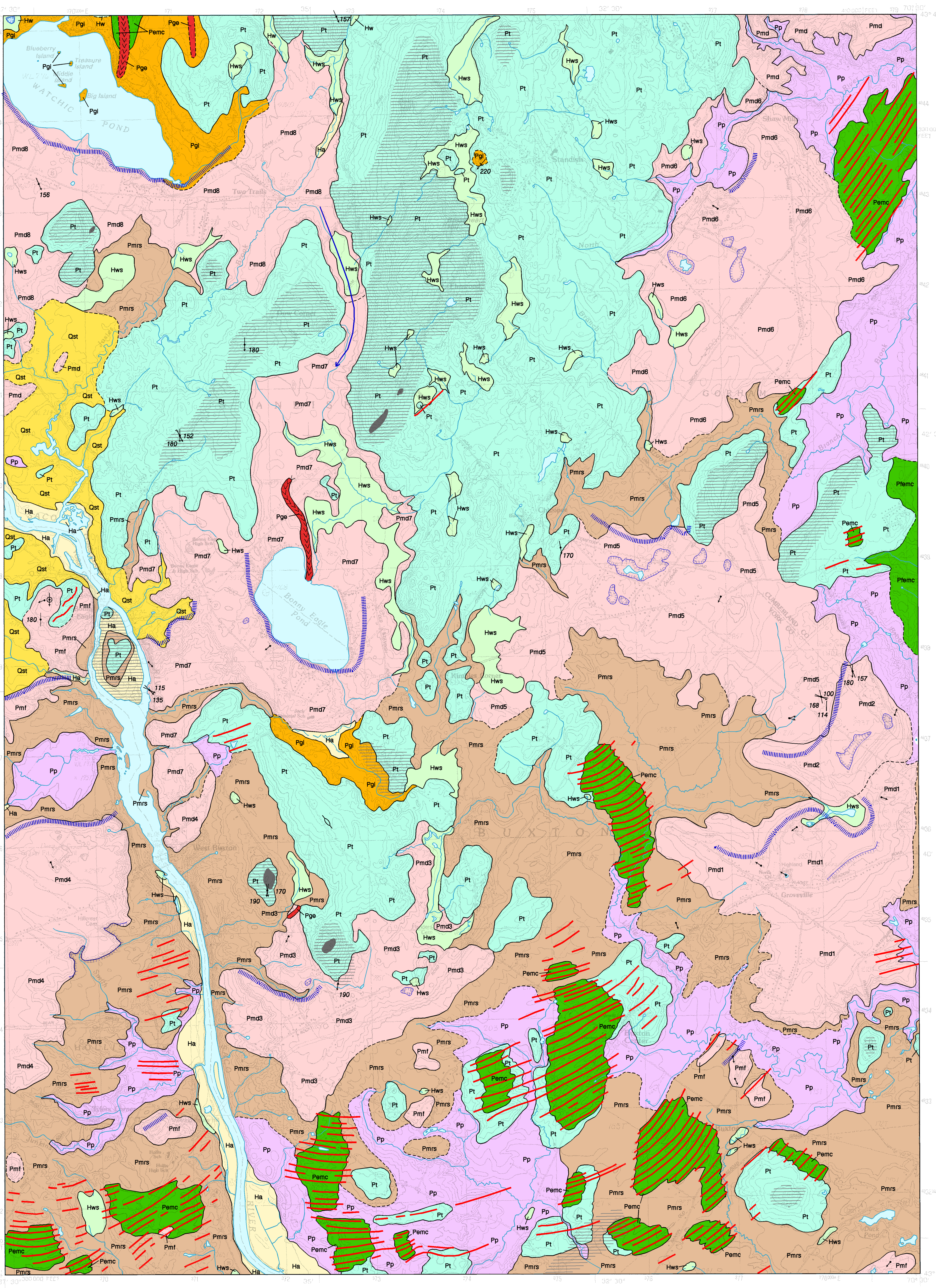
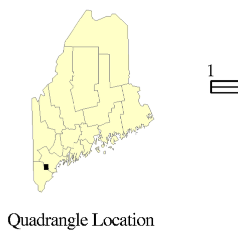
Surficial Geology



SOURCES OF INFORMATION

Surficial geologic mapping by John C. Gosse completed during the 1990 field season; funding for this work provided by the U. S. Geological Survey COGEOMAP program. Geologic unit designations and contacts revised and matched to adjacent quadrangles in 1999 by MGS geologists.



Quadrangle Location

SCALE 1 : 24,000

1 — 0 — 1 MILE

1000 0 1000 2000 3000 4000 5000 6000 7000 FEET

1 — 0 — 1 KILOMETER

CONTOUR INTERVAL 10 FEET

TRUE NORTH

Topographic base from U.S. Geological Survey Standish quadrangle, scale 1:24,000 using standard U.S. Geological Survey topographic map symbols.

The use of industry, firm, or local government names on this map is for location purposes only and does not impute responsibility for any present or potential effects on the natural resources.

Ha — **Alluvium** - Generally well-sorted and stratified silt, sand, and gravel on flood plains of modern rivers and streams.

Hw — **Freshwater wetlands** - Muck, peat, silt, and sand. Poorly drained areas, often with standing water.

Hws — **Wetland, swamp** - Peat, muck, silt, and clay in poorly drained and wetland areas with variable tree cover. Includes marshes and heaths.

Qst — **Stream terrace** - Generally flat surface planed by postglacial fluvial activity or underlain by alluvium. Surface is above modern flood plain. Associated with postglacial river and stream deposits.

Pmrs — **Marine regressive sand deposits** - Well-sorted massive to finely laminated silt to coarse sand. Associated with Pp, Pmd, and Pmf, which it commonly overlies. Locally includes shoreline and nearshore deposits.

Pp — **Presumpscot Formation** - Greenish-gray to bluish-gray marine silt and clay. May be laminated or massive. Occurs as variably thick veneer overlying older units below the marine limit.

Pmd — **Marine delta** - Generally thick deposits of sorted and interbedded sand and gravel with well-developed deltaic foreset beds. Topset beds are known or inferred to be present. Grades seaward to Pmrs or Pp. Locally includes shoreline and nearshore sand and gravel deposits where delta was eroded by the sea. If present, a subscript number refers to sequence of deposition (Table 2 of Gosse, 1999).

Pmf — **Marine fan** - Generally thick deposits of sorted and interbedded sand and gravel. Foreset beds commonly well-developed. No topset beds. Grades seaward to Pmrs or Pp. Closely associated with Pemc.

Pge — **Esker** - Sinuous, generally discontinuous ridges of massive and stratified, commonly interbedded, sand and gravel which was released from subglacial and englacial meltwater conduits during glacial retreat. Cobble to boulder-size clasts are commonly highly rounded and spherical. Associated with Pmd.

Pgi — **Ice-contact deposits** - Massive to stratified seaward-graded sand and gravel deposited near the ice margin by glacial meltwater. May include outwash. Topography locally kettled or hummocky. Evidence of deformation (faults and folds) often present.

Pfemc — **Fan-end moraine complex** - Composite unit incorporating elements of end moraines and subaqueous fans. Coarse to fine sand and gravel extending from fan head to fan toe. This material overlies sediments of end moraine and end moraine complex.

Pemc — **End moraine complex** - Cluster of closely (and often evenly) spaced ridges comprised of till or poorly sorted sand and gravel, deposited at the glacier margin. Can include small patches of Pmf and thin veneers of Pp or Pmrs.

Pt — **Till** - Homogeneous, locally compact, poorly sorted mixture of a wide range of particle sizes from silt to boulders. Unit is widespread with very variable thickness (typically 0 - 5 m). Associated with bedrock highs.

Bedrock - Bedrock of Paleozoic age. "rk" indicates areas of barren ledge. The ruled pattern indicates areas where surficial sediments are generally less than 2 m thick.

Contact - Boundary between map units. Dashed where boundary is uncertain or inferred.

Ice margin position - Inferred position of the glacier margin during deposition of the adjacent marine delta or fan. The line is drawn along the backslope of deltas.

Scarp - Wave-cut scarp eroded at front of delta during regression of sea.

Moraine ridge - Ridges of till and/or water-laid sediments deposited in the marginal zone of the glacier. Lines indicate ridge crests.

>>>>> **Esker ridge** - Shows trend of sand and gravel ridge deposited in meltwater tunnel beneath the ice sheet (Pge). Chevrons indicate the direction of inferred meltwater flow.

135 **Glacial striation locality** - Dot indicates point of observation. Arrow indicates direction of ice flow, if known. Flagged arrows indicate earlier flows (see Table 3 of Gosse, 1999).

135 **Till fabric locality** - Combined orientation of the long axis of at least 30 pebbles in till (see Figure 3 of Gosse, 1999). Dot indicates fabric site. Arrow indicates direction of inferred ice flow.

135 **Fold axis** - Measurement of the trend of a fold axis in deformed ice-contact sediment. Intersection of the bar and arrow indicates the location of the observation. Arrow indicates the inferred ice flow direction (see Table 3 of Gosse, 1999).

Paleoflow direction - Arrow indicates the flow direction of a paleocurrent which deposited glaciofluvial or glaciomarine sediments. Tail of arrow indicates location of observation (see Table 4 of Gosse, 1999).

Glacially streamlined hill - Indicates a hill that has been elongated parallel to the direction of ice flow. The hill may be bedrock-cored.

Kettle - Depression created by melting of a buried mass of glacial ice and collapsing of the overlying sediment. May contain a small pond or wetland.

Meltwater channel - Channel cut by glacial meltwater stream.

USES OF SURFICIAL GEOLOGY MAPS

A surficial geology map shows all the loose materials such as till (commonly called hardpan), sand and gravel, or clay, which overlie solid ledge (bedrock). Bedrock outcrops and areas of abundant bedrock outcrops are shown on the map, but varieties of the bedrock are not distinguished (refer to bedrock geology map). Most of the surficial materials are deposits formed by glacial and deglacial processes during the last stage of continental glaciation, which began about 25,000 years ago. The remainder of the surficial deposits are the products of postglacial geologic processes, such as river floodplains, or are attributed to human activity, such as fill or other land-modifying features.

The map shows the areal distribution of the different types of glacial features, deposits, and landforms as described in the map explanation. Features such as striations and moraines can be used to reconstruct the movement and position of the glacier and its margin, especially as the ice sheet melted. Other ancient features include shorelines and deposits of glacial lakes or the glacial sea, now long gone from the state. This glacial geologic history of the quadrangle is useful to the larger understanding of past earth climate, and how our region of the world underwent recent geologically significant climatic and environmental changes. We may then be able to use this knowledge in anticipation of future similar changes for long-term planning efforts, such as coastal development or waste disposal.

Surficial geology maps are often best used in conjunction with related maps such as surficial materials maps or significant sand and gravel aquifer maps for anyone wanting to know what lies beneath the land surface. For example, these maps may aid in the search for water supplies, or economically important deposits such as sand and gravel for aggregate or clay for bricks or pottery. Environmental issues such as the location of a suitable landfill site or the possible spread of contaminants are directly related to surficial geology. Construction projects such as locating new roads, excavating foundations, or siting new homes may be better planned with a good knowledge of the surficial geology of the site. Refer to the list of related publications below.

OTHER SOURCES OF INFORMATION

1. Gosse, J. C., 1999, Surficial geology of the Standish 7.5-minute quadrangle, York and Cumberland Counties, Maine: Maine Geological Survey, Open-File Report 99-132, 24 p.
2. Gosse, J. C., 1998, Surficial materials of the Standish quadrangle, Maine: Maine Geological Survey, Open-File Map 98-176.
3. Neil, C. D., 1998, Significant sand and gravel aquifers of the Standish quadrangle, Maine: Maine Geological Survey, Open-File Map 98-142.
4. Thompson, W. B., 1979, Surficial geology handbook for coastal Maine: Maine Geological Survey, 68 p. (out of print)
5. Thompson, W. B., and Borns, H. W., Jr., 1985, Surficial geologic map of Maine: Maine Geological Survey, scale 1:500,000.
6. Thompson, W. B., Crossen, K. J., Borns, H. W., Jr., and Andersen, B. G., 1989, Glaciomarine deltas of Maine and their relation to late Pleistocene-Holocene crustal movements, *in* Anderson, W. A., and Borns, H. W., Jr. (eds.), Neotectonics of Maine: Maine Geological Survey, Bulletin 40, p. 43-67.

Standish Quadrangle, Maine

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1999

For additional information, see Open-File Report 99-132.

SURFICIAL GEOLOGY OF MAINE

Continental glaciers like the ice sheet now covering Antarctica probably extended across Maine several times during the Pleistocene Epoch, between about 1.5 million and 10,000 years ago. The slow-moving ice superficially changed the landscape as it scraped over mountains and valleys (**Figure 1**), eroding and transporting boulders and other rock debris for miles (**Figure 2**). The sediments that cover much of Maine are largely the product of glaciation. Glacial ice deposited some of these materials, while others washed into the sea or accumulated in meltwater streams and lakes as the ice receded. Earlier stream patterns were disrupted, creating hundreds of ponds and lakes across the state. The map at left shows the pattern of glacial sediments in the Standish quadrangle.

The most recent "Ice Age" in Maine began about 25,000 years ago, when an ice sheet spread southward over New England (Stone and Borns, 1986). During its peak, the ice was several thousand feet thick and covered the highest mountains in the state. The weight of this huge glacier actually caused the land surface to sink hundreds of feet. Rock debris frozen into the base of the glacier abraded the bedrock surface over which the ice flowed. The grooves and fine scratches (striations) resulting from this scraping process are often seen on freshly exposed bedrock, and they are important indicators of the direction of ice movement (**Figure 3**). Erosion and sediment deposition by the ice sheet combined to give a streamlined shape to many hills, with their long dimension parallel to the direction of ice flow. Some of these hills (drumlins) are composed of dense glacial sediment (till) plastered under great pressure beneath the ice.

A warming climate forced the ice sheet to start receding as early as 21,000 years ago, soon after it reached its southernmost position on Long Island (Sirkin, 1986). The edge of the glacier withdrew from the continental shelf east of Long Island and reached the present position of the Maine coast by 13,800 years ago (Dorion, 1993). Even though the weight of the ice was removed from the land surface, the Earth's crust did not immediately spring back to its normal level. As a result, the sea flooded much of southern Maine as the glacier retreated to the northwest. Ocean waters extended far up the Kennebec and Penobscot valleys, reaching present elevations of up to 420 feet in the central part of the state.

Great quantities of sediment washed out of the melting ice and into the sea, which was in contact with the receding glacier margin. Sand and gravel accumulated as deltas (**Figure 4**) and submarine fans where streams discharged along the ice front, while the finer silt and clay dispersed across the ocean floor. The shells of clams, mussels, and other invertebrates are found in the glacial-marine clay that blankets lowland areas of southern Maine. Age dates on these fossils tell us that ocean waters covered parts of Maine until about 11,000 years ago, when the land surface rebounded as the weight of the ice sheet was removed.

Meltwater streams deposited sand and gravel in tunnels within the ice. These deposits remained as ridges (eskers) when the surrounding ice disappeared (**Figure 5**). Maine's esker systems can be traced for up to 100 miles, and are among the longest in the country.

Other sand and gravel deposits formed as mounds (kames) and terraces adjacent to melting ice, or as outwash in valleys in front of the glacier. Many of these water-laid deposits are well layered, in contrast to the chaotic mixture of boulders and sediment of all sizes (till) that was released from dirty ice without subsequent reworking. Ridges consisting of till or washed sediments (moraines) were constructed along the ice margin in places where the glacier was still actively flowing and conveying rock debris to its terminus. Moraine ridges are abundant in the zone of former marine submergence, where they are useful indicators of the pattern of ice retreat (**Figure 6**).

The last remnants of glacial ice probably were gone from Maine by 10,000 years ago. Large sand dunes accumulated in late-glacial time as winds picked up outwash sand and blew it onto the east sides of river valleys, such as the Androscoggin and Saco valleys (**Figure 7**). The modern stream network became established soon after deglaciation, and organic deposits began to form in peat bogs, marshes, and swamps. Tundra vegetation bordering the ice sheet was replaced by changing forest communities as the climate warmed (Davis and Jacobson, 1985). Geologic processes are by no means dormant today, however, since rivers and wave action modify the land (**Figure 8**), and worldwide sea level is gradually rising against Maine's coast.

References Cited

Davis, R. B., and Jacobson, G. L., Jr., 1985, Late-glacial and early Holocene landscapes in northern New England and adjacent areas of Canada: Quaternary Research, v. 23, p. 341-368.

Dorion, C. C., 1993, A chronology of deglaciation and accompanying marine transgression in Maine: Geological Society of America, Abstracts with Programs, v. 25, no. 2, p. 12.

Sirkin, L., 1986, Pleistocene stratigraphy of Long Island, New York, *in* Caldwell, D. W. (editor), The Wisconsinan stage of the first geological district, eastern New York: New York State Museum, Bull. 455, p. 6-21.

Stone, B. D., and Borns, H. W., Jr., 1986, Pleistocene glacial and interglacial stratigraphy of New England, Long Island, and adjacent Georges Bank and Gulf of Maine, *in* Sibrava, V., Bowen, D. Q., and Richmond, G. M. (editors), Quaternary glaciations in the northern hemisphere: Quaternary Science Reviews, v. 5, p. 39-52.

Figure 1: "The Bubbles" and Jordan Pond in Acadia National Park. These hills and valleys were sculpted by glacial erosion. The pond was dammed behind a moraine ridge during retreat of the ice sheet.

Figure 2: Daggett's Rock in Phillips. This is the largest known glacially transported boulder in Maine. It is about 100 feet long and estimated to weigh 8,000 tons.

Figure 3: Granite ledge in Westbrook, showing polished and grooved surface resulting from glacial abrasion. The grooves and shape of the ledge indicate ice flow toward the southeast.

Figure 4: Glaciomarine delta in Franklin, formed by sand and gravel washing into the ocean from the glacier margin. The flat delta top marks approximate former sea level. Kettle hole in foreground was left by melting of ice.

Figure 5: Esker cutting across Kezar Five Ponds, Waterford. The ridge consists of sand and gravel deposited by meltwater flowing in a tunnel beneath glacial ice.

Figure 6: Aerial view of moraine ridges in blueberry field, Sedgwick (note dirt road in upper right for scale). Each bouldery ridge marks a position of the retreating glacier margin. The ice receded from right to left.

Figure 7: Sand dune in Wayne. This and other "deserts" in Maine formed as windstorms in late-glacial time blew sand out of valleys, often depositing it as dune fields on hillsides downwind. Some dunes were reactivated in historical time when grazing animals stripped the vegetation cover.

Figure 8 : Songo River delta and Songo Beach, Sebago Lake State Park, Naples. These deposits are typical of geological features formed in Maine since the Ice Age.